

Anthropogenic carbon and ocean pH

Ken Caldeira^{1*} and Michael E. Wickett²

¹*Energy and Environment Directorate, Lawrence Livermore National Laboratory,
7000 East Ave, Livermore, CA 94550, USA*

²*Center for Applied Scientific Computing, Lawrence Livermore National Laboratory,
7000 East Ave, Livermore, CA 94550, USA*

*To whom correspondence should be addressed; E-mail: kenc@llnl.gov

Most fossil-fuel CO₂ released to the atmosphere will eventually be absorbed by the ocean¹ with potentially adverse consequences for marine biota²⁻⁴. We quantify pH changes that may result from continued release of fossil-fuel CO₂ to the atmosphere, and compare these with pH changes estimated from geological and historical records. We conclude that releasing fossil-fuel CO₂ to the atmosphere over several centuries may result in ocean pH changes greater than any inferred from the geologic record of the past 300 million years, with the possible exception of rare extreme events such as bolide impacts or catastrophic methane hydrate degassing.

When carbon dioxide dissolves in the ocean it lowers the pH, making the ocean more acidic. Due to a paucity of relevant observations, we have a limited understanding of the effects of pH reductions on marine biota. Coral reefs², calcareous plankton³, and other organisms with calcium carbonate skeletal material or shells may be particularly impacted. Most biota resides near the surface where greatest pH change can be expected to occur, but deep ocean biota may be more sensitive to pH changes⁴.

To study effects of CO₂ emissions on ocean pH, we forced the LLNL ocean general circulation model⁵ (Fig. 1a) with observed atmospheric pCO₂ from 1975 to 2000 and with CO₂ emissions from the IPCC IS92a scenario¹ from year 2000 to 2100. Beyond 2100, emissions follow a logistic function burning remaining fossil fuel resources (assuming 5270 GtC in year 1750; ref. 6, 7). Simulated atmospheric CO₂ exceeds 1900 ppm near

year 2300. Maximum surface ocean pH decrease is 0.77 . We estimate, using the geochemical model^{8,9} described below, that changes in temperature, weathering, and sedimentation would reduce this maximum decrease by <10%.

A review¹⁰ of estimates of paleo-atmospheric CO₂ from geochemical models, paleosols, algae and forams, plant stomata, and boron isotopes concluded that there is no evidence that concentrations exceeded 7,500 ppm or were less than 100 ppm over the past 300 myr. In addition, the highest concentrations inferred from the geologic record were thought to have developed over many millions of years due to slow processes involving tectonics and biological evolution.

We estimated the effect of past changes in atmospheric CO₂ on ocean pH with a four-box ocean/atmosphere model^{8,9}. Modeled processes include weathering of carbonate and silicate minerals on land, production of shallow-water carbonate minerals, production and oxidation of biogenic organic carbon, production and dissolution of biogenic carbonate minerals in the ocean, air-sea gas exchange of carbon, and transport by advection, mixing, and biology.

In a series of simulations, atmospheric pCO₂ was varied linearly from the pre-industrial value (~280 ppm) to stabilization values ranging from 100 ppm to 10,000 ppm over time intervals ranging from 10 yr to 10⁷ yr. For each model simulation, we recorded maximum predicted pH perturbation in the surface ocean boxes (Fig. 1b). When a CO₂ change occurs over a short time interval (i.e., <~10⁴ yr), ocean pH is relatively sensitive to added CO₂. However, when a CO₂ change occurs over a long time interval (i.e., >~10⁵ yr), ocean chemistry is buffered by interactions with carbonate minerals, lowering pH sensitivity¹¹.

Based on the record¹⁰ of atmospheric CO₂ over the past 300 myr and our geochemical model^{8,9}, there is no evidence that ocean pH was more than 0.6 units lower than today. Our GCM results indicate that continued fossil-fuel burning with atmospheric CO₂ release could lead to pH decreases of 0.7 units. Thus, we conclude that unabated

CO₂ emissions over the next several hundred years may produce changes in ocean pH greater in magnitude than any experienced in the past 300 myr, with the possible exception of rare catastrophic events in Earth history^{8,12}.

References

1. Houghton, J.T. et al. (eds), *Climate Change 2001 - the scientific basis. Contribution of Working Group I to the Third Assessment Report of the Intergovernmental Panel on Climate Change*. Cambridge University Press, 944 pp, (2001).
2. Kleypas, J. A. et al. Geochemical consequences of increased atmospheric carbon dioxide on coral reefs. *Science* **284**, 118–120 (1999).
3. Riebesell, U. et al. Reduced calcification of marine plankton in response to increased atmospheric CO₂. *Nature* **407**, 364–367 (2000).
4. Seibel, B. A. & Walsh, P. J. Carbon cycle — Potential impacts of CO₂ injection on deep-sea biota. *Science* **294**, 319–320 (2001).
5. Caldeira, K. & Duffy, P. B. The role of the Southern Ocean in uptake and storage of anthropogenic carbon dioxide. *Science* **287**, 620–622 (2000).
6. Metz, B, Davidson, O., Swart, R. & Pan, J. (Eds.) *Climate Change 2001 - mitigation. Contribution of Working Group III to the Third Assessment Report of the Intergovernmental Panel on Climate Change*. Cambridge University Press, 700 pp. (2001).
7. Marland, G., Boden, T.A. & Andres, R.J. Global, Regional, and National Fossil Fuel CO₂ Emissions. In *Trends: A Compendium of Data on Global Change*. Carbon Dioxide Information Analysis Center, Oak Ridge National Laboratory, U.S. Department of Energy, Oak Ridge, Tenn., U.S.A (2002).

8. Caldeira, K. & Rampino, M. R. Aftermath of the end-Cretaceous mass extinction – possible biogeochemical stabilization of the carbon cycle and climate. *Paleoceanography* **8**, 515–525 (1993).
9. Caldeira, K. & Rau, G. H. Accelerating carbonate dissolution to sequester carbon dioxide in the ocean: Geochemical implications. *Geophys. Res. Letters* **27**, 225–228 (2000).
10. Crowley, T. J. & Berner, R. A. Paleoclimate — CO₂ and climate change. *Science* **292**, 780–781 (2001).
11. Caldeira, K. & Berner, R. Seawater pH and atmospheric carbon dioxide (Technical comment) *Science* **286**, 2043a (1999).
12. Beerling, D. J. & Berner, R. A. Biogeochemical constraints on the Triassic-Jurassic boundary carbon cycle event. *Global Biogeochem. Cycles* **16**, 101–113 (2002).
13. Sanyal, A., Hemming, N. G., Hanson, G. N. & Broecker, W. S. Evidence for a higher pH in the glacial ocean from boron isotopes in foraminifera. *Nature* **373**, 234–236 (1995).

Acknowledgements This research was supported by the DOE Office of Science Office of Biological and Environmental Research Ocean Carbon Sequestration Research Program. This work was performed under the auspices of the U.S. Department of Energy by University of California Lawrence Livermore National Laboratory under contract No. W-7405-Eng-48.

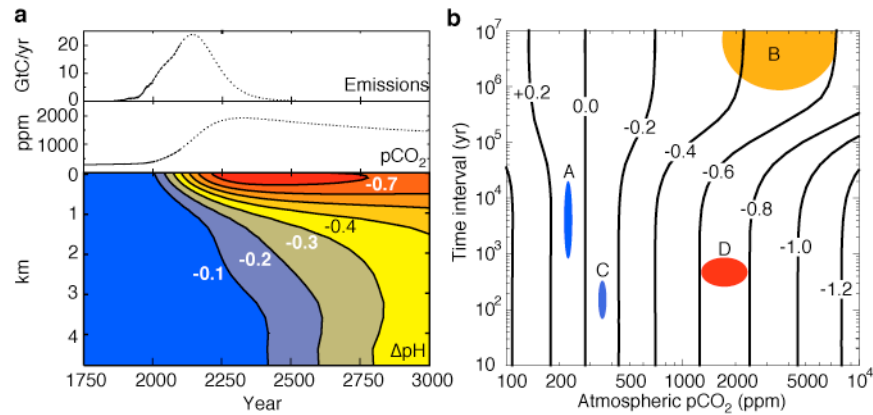


Figure 1 Atmospheric release of fossil fuel CO₂ may make the ocean more acid than it has been for the past 300 myr. **a**, Atmospheric CO₂ emissions, historical atmospheric CO₂ and CO₂ predicted from this emissions scenario, and change in ocean pH based on horizontally averaged chemistry. **b**, Estimated maximum surface ocean pH change as a function of final atmospheric pCO₂ and the transition time over which this pCO₂ is linearly approached from 280 ppm: A, glacial-interglacial CO₂ changes¹³; B, slow changes over the past 300 myr; C, historical changes¹ in ocean surface waters; D, unabated fossil-fuel burning over the next few centuries.