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Atmospheric CO₂ and average ocean Δ pH

How sensitive are ocean systems to pH changes? If we permit atmospheric CO₂ levels to rise to ~ 600 ppm and remain there, then surface ocean waters, and eventually all ocean waters, will experience an ~ 0.3 pH change from pre-industrial levels. Is that an acceptable limit, and why? What burden of fossil fuel CO₂ is "acceptable" for climate, or geochemical, or biological, reasons? How might we view the trade-off between the direct effects of CO₂ and of the induced effects of climate/temperature? The focus of the ocean science community has been on observing the evolving fossil fuel CO₂ tracer signal, and the questions above have simply (and astonishingly) not yet been posed.

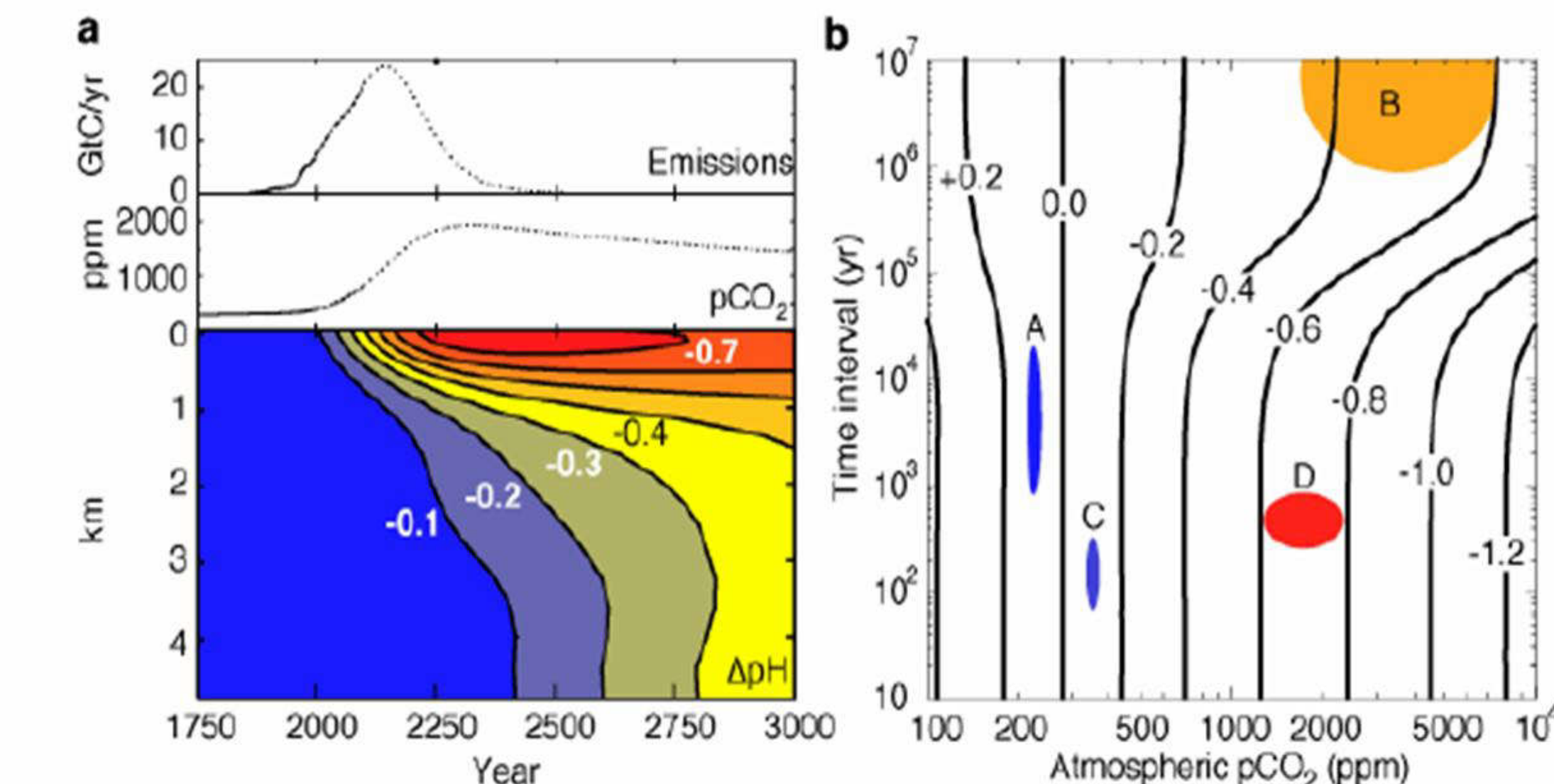
[1] Brewer, P.G. (1997) Ocean chemistry of the fossil fuel CO₂ signal: The haline signature of "Business as Usual". *Geophys. Res. Lett.* 24 1367-1369.

[2] Langdon, C., T. Takahashi, C. Sweeney, D. Chipman, J. Goddard, F. Marubini, H. Aceves, H. Barnett, and M.J. Atkinson (2000) Effect of calcium carbonate saturation state on the calcification rate of an experimental coral reef. *Global Biogeochem. Cycles* 14 639-654.

[3] McNeil, B., R.J. Matear, R.M. Key, J.L. Bullister, and J. Sarmiento. (2003) Anthropogenic CO₂ uptake by the ocean based on the global chlorofluorocarbon data set. *Science*, 299, 235-239.

[4] Riebesell, U., I. Zondervan, B. Rost, P.D. Tortell, R.E. Zeebe, and F. Moll (2000) Reduced calcification of marine plankton in response to increased atmospheric CO₂. *Nature*, 407, 364-367.

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Atmospheric release of fossil fuel CO₂ may make the ocean more acid than it has been for the past 300 myr. a, Atmospheric CO₂ emissions, historical atmospheric CO₂ and CO₂ predicted from this emissions scenario, and change in ocean pH based on horizontally averaged chemistry. b, Estimated maximum surface ocean pH change as a function of final atmospheric pCO₂ and the transition time over which this pCO₂ linearly approached from 280 ppm: A, glacial-interglacial CO₂ changes¹³; B, slow changes over the past 300 myr; C, historical changes¹ in ocean surface waters; D, unabated fossil-fuel burning over the next few centuries. ^[5]

CO2 injection causes;

- 1 – Elevated CO2 levels
- 2 – Elevated Seawater Acidity (reduced pH)

Respiratory Stress
Increased acidity reduces the level of oxygen-binding and transport by respiratory proteins, leading to reduced aerobic capacity.

Acidosis (reduced internal pH)
Disruption of the acid/base balance of animal tissues impairs function and requires energy to restore or maintain optimal internal pH levels.

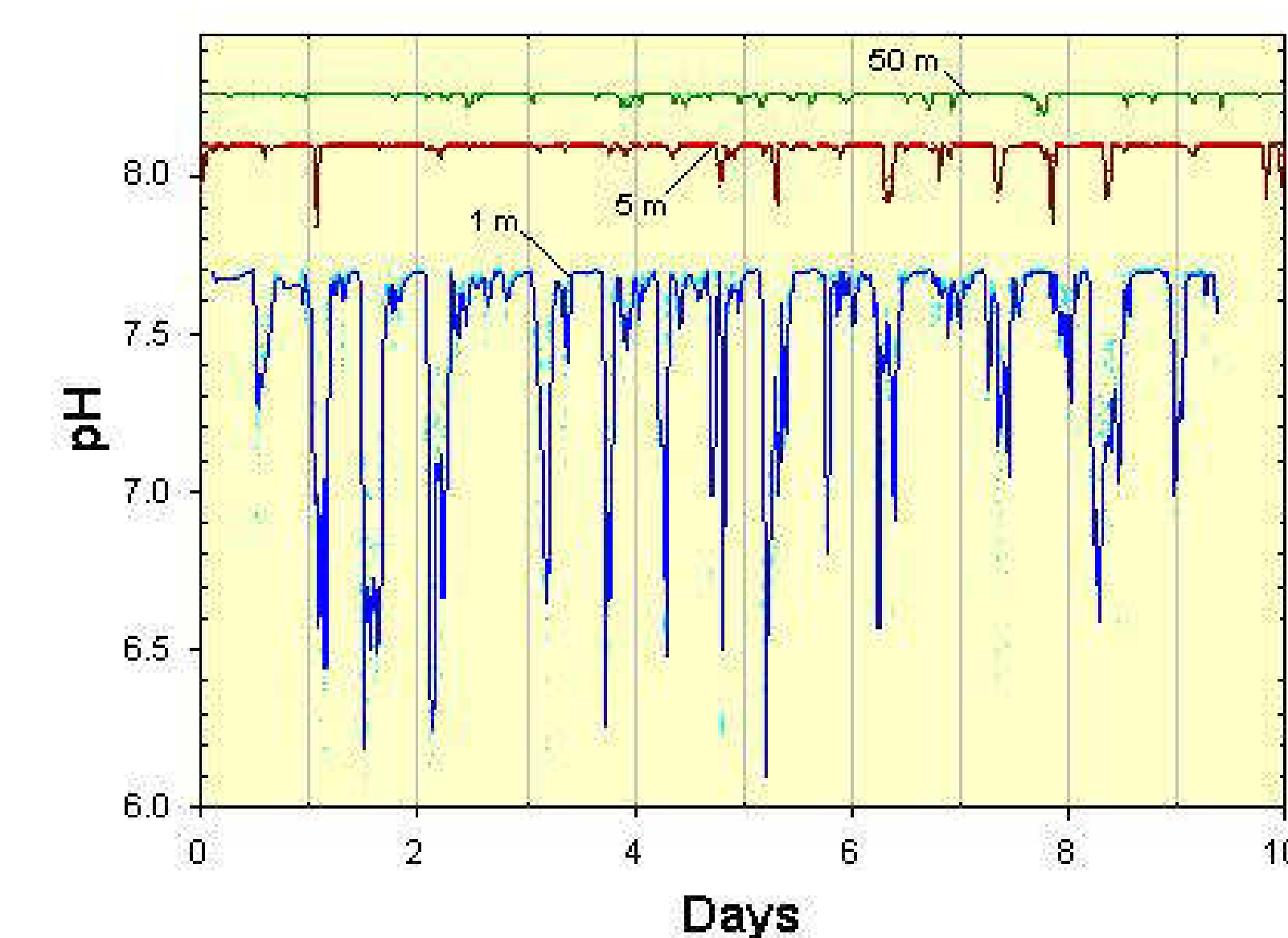
Metabolic suppression (Torpor)
Elevated CO₂, reduced pH, or both can cause some animals to enter a state of reduced metabolic rate and semi-hibernation, reducing their protein synthesis, growth, reproduction, etc.

- Degraded

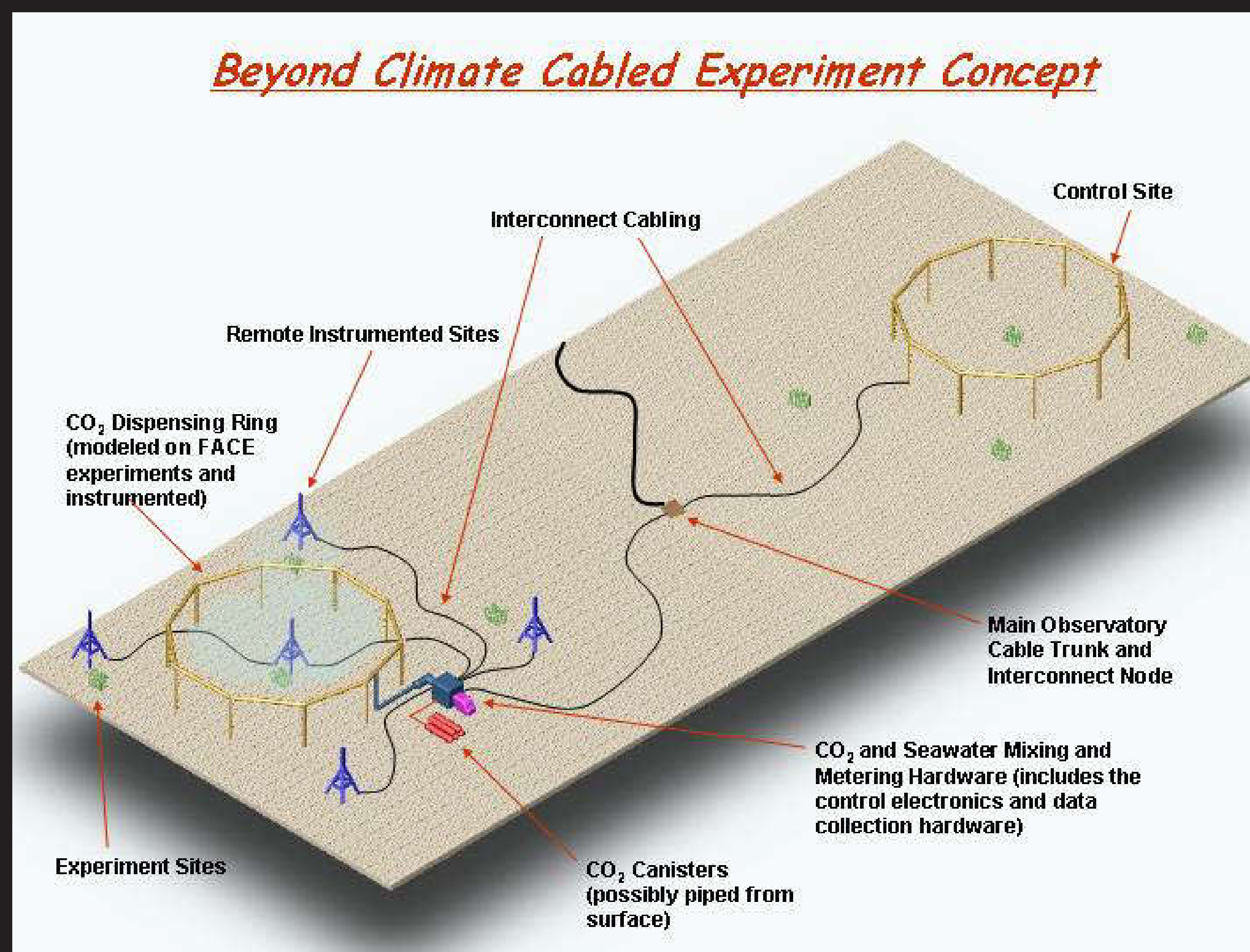
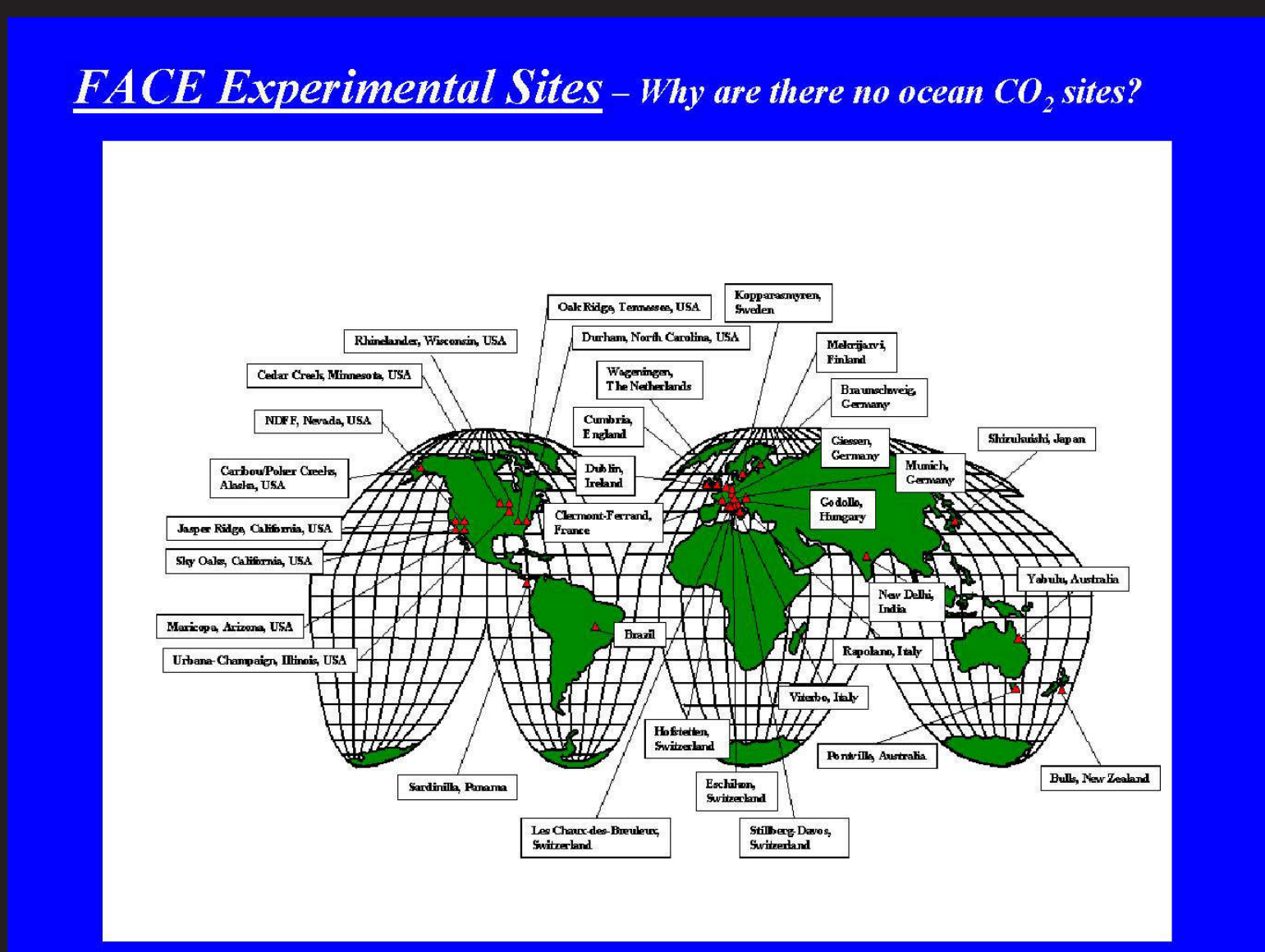


We can view the current series of deep ocean CO₂ injection experiments not simply as disposal strategy tests, but also as controlled enrichment experiments that may allow us to mimic the elevated CO₂ levels of a future ocean in much the same way that ecosystems are manipulated on land (Shaw et al., 2002; DeLucia et al., 1999). While these oceanic injection experiments have broken new ground, they are deficient in many ways: it is hard to control the pH perturbation achieved due to the varying tidal forcing (See Figure on right), and very close to the source the CO₂ gradients are so large that animals may either not survive, or show little response, depending upon very small changes in position. The experiments carried out so far have been short term due to limits on the amount of CO₂ that can be deployed by an ROV.

Spiking due to currents and tides causing near field animals and experimental results to be questionable over time. Beyond CIn experiment will address these inaccuracies by regulating the fluctuation accurately represent a future pH ocean change



Face Experiment Structure as Background



Designing, testing, and executing a sophisticated set of controlled perturbation experiments on oceanic biogeochemical systems will involve the creation of a large number of scientists and engineers. While we can learn from what has been carried out on land, the translation to oceanic systems is not easy. Among the problems to be considered are:

- A) On land the trees, shrubs, grasses etc. are immobile - oceanic animals move, and the protocols for obtaining valid results will have to be carefully evaluated.
- B) CO₂ has no atmospheric chemistry, and mixing the source gas into ambient air is a simple problem. Adding CO₂ to ocean waters involves the full set of aqueous CO₂ system changes, and the lag times associated with them.
- C) At depths typically > -350m CO₂ will react with sea water to form a hydrate, and flow assurance will have to be planned for.
- D) The time scale for observing changes may be quite long, for example in the case of carbonate sediment dissolution experiments.
- E) The challenge of observing the perturbed area must be addressed by novel sensor techniques - simply taking a large number of cores may damage the experimental site.

However none of the problems are insoluble, and much is known already. In time control and data flow very rapid advances can be made.

A) It is possible to extend the work beyond CO₂, for example to perturbations of the oxygen content, or temperature, of the system.

B) A wide range of pH perturbations may be studied.

C) The earliest experimental targets should be in shallower waters, including coral reef systems.

This will permit rapid system development, and provide guidance for steadily deeper deployments.